

New Eyes in the Sky Measure Glaciers and Ice Sheets

The mapping and measurement of glaciers and their changes are useful in predicting sea-level and regional water supply, studying hazards and climate change [Haerberli *et al.*, 1998], and in the hydropower industry. Existing inventories cover only about 67,000 of the world's estimated 160,000 glaciers and are based on data collected over 50 years or more [e.g., Haerberli *et al.*, 1998]. The data available have proven that small ice bodies are disappearing at an accelerating rate and that the Antarctic ice sheet and its fringing ice shelves are undergoing unexpected, rapid change. According to many glaciologists, much larger fluctuations in land ice—with vast implications for society—are possible in the coming decades and centuries due to natural and anthropogenic climate change [Oppenheimer, 1998].

Before glacier data can be used to address critical problems pertinent to the world's economic and environmental health, the data need to become more comprehensive, more homogenous in detail and quality, and managed within Geographic Information Systems (GIS). A synchronous, GIS-based, uniform baseline inventory needs to be established to track global glacier changes; ground-validated satellite measurements are crucial to extend coverage to a representative sample of the world's land ice [Dyrugorov and Meier, 1997a,b; Haerberli, 1998].

The launch by NASA of two new environmental satellites (Landsat-7 and Terra) in 1999 marks a leap in capabilities for global mapping and measurements of glaciers and ice sheets. An international glaciological consortium, Global Land Ice Measurements from Space (GLIMS), has been funded by NASA as a "Pathfinder" project to assess the status and vigor of the world's glaciers using remote sensing and field data (<http://www.flag.wr.usgs.gov/GLIMS/>). GLIMS and similar projects are using satellite imaging to provide our first comprehensive and uniform look at the world's land ice.

New Imaging Platforms

The Landsat-7 global glacier observation project is already quickly developing. This satellite's Enhanced Thematic Mapper Plus (ETM+, <http://landsat.gsfc.nasa.gov/>) is fully operational and delivers a veritable flood of images of glaciers and ice sheets (Figure 1). The L-7 data acquisition plan calls for three image acquisition attempts each summer toward the end of the melt season (spaced

every 16-day ground-track repeat cycle) for all temperate glaciers, four acquisitions per year (once every three months) of Andean and equatorial glaciers, and one acquisition in mid-summer of Greenland and Antarctica. Already, L-7 coverage exists for most of Antarctica and Greenland and many alpine glacier regions. The L-7 plan also has many spatial-temporal blocks of sea-ice collections.

A Japanese instrument, the Advanced Spaceborne Thermal Emission and Reflection Radiometer (ASTER), (<http://asterweb.jpl.nasa.gov/aster.pdf>), was launched by NASA aboard Terra on December 18, 1999, and is expected to begin delivering an equivalent flood of multispectral, stereo imaging of glaciers starting in the Northern Hemisphere in June 2000, and in the Southern Hemisphere in December 2000. Test images received to date are of high quality. GLIMS has an

approach to ASTER imaging that, somewhat similar to the plan for Landsat 7, involves repeated imaging attempts late in the melt season. A system for making semi-automated glacier image data acquisition requests for ASTER was recently described by Raup *et al.* (2000).

Terra is orbiting in formation with Landsat-7, so illumination geometry of ASTER and ETM+ images is virtually the same. ASTER and ETM+ have complementary capabilities. ETM+ has a wide swath width (185 km) and is thus better suited than ASTER to cover the large expanses of polar ice sheets. ASTER, with a 60-km swath width and stereo-imaging capability, is better suited to cover alpine glaciers, which tend to be debris-covered at their termini and are best mapped using stereo images. ASTER and ETM+ are similar in terms of wavelength coverage, except that ASTER has more thermal imaging bands. With 15-m spatial resolution in three VNIR bands (including one band in stereo), 30-m resolution in six SWIR bands, and 90-m resolution in five TIR bands, ASTER is well-equipped to study the fine structural details of small alpine glaciers and to map snow, ice, and debris-covered glacier facies and moraines.

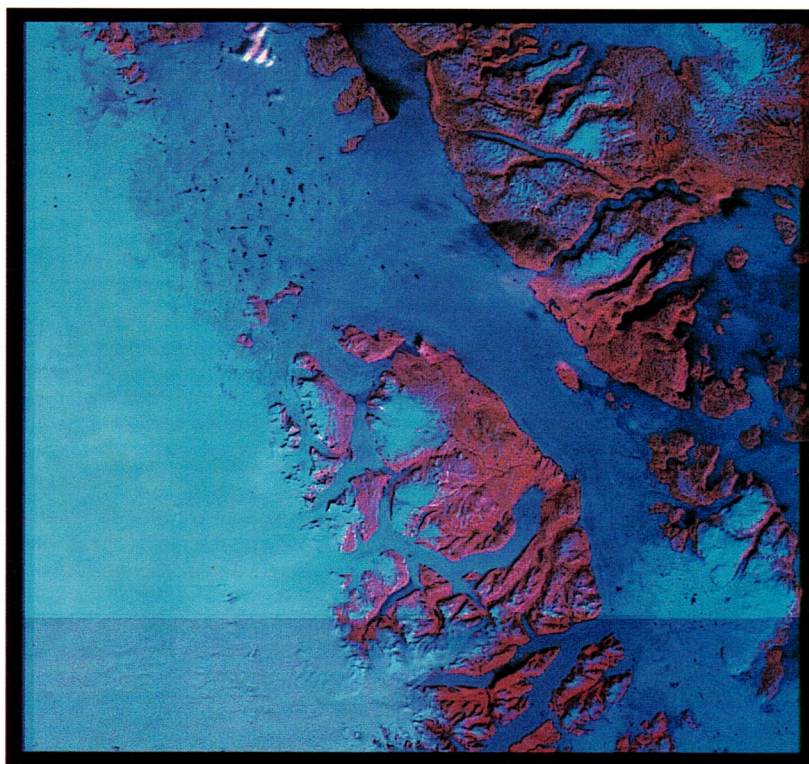


Fig. 1. Landsat-7 image of Storstrommen glacier, northeast Greenland, in 1999 showing many icebergs at the calving front. The front of Storstrommen surged by more than 10 km [Higgins, 1994] between 1978 and 1984. During this period, its average surface velocity increased from a normal several hundred meters per year to as much as 4000 m/yr. Transfer of ice from the upper to the lower part of the ablation area during this surge resulted in an 80-m lowering of the glacier surface in the upper part of the ablation area and a rise of similar magnitude in the lower part. Landsat-7 and ASTER images will be used to continue the study of this and other glaciers of Greenland.

An example of this type of glacier image analysis is shown in Figure 3. The region shown, Nanga Parbat (Himalaya), is especially difficult to deal with using multispectral remote sensing methods because of the extreme relief and heavy debris cover of glaciers [Bishop *et al.*, 1998]. The chosen method of glacier mapping requires development of a relative topographic Digital Terrain Model (DTM), which in principle could be produced from existing topographic maps or from stereo images. DTM algorithms were written at USGS so as to provide an inexpensive and reliable alternative to existing maps and commercial DTM software. Since DTMs are rarely available at the resolution needed, and since the maps are often old and glaciers are changing, ASTER stereo imaging is expected to provide

the primary data source for stereophotogrammetric DTMs. Existing topographic maps covering selected glaciers will provide a check on topography extracted from satellite data. Older topographic maps will provide information on glacier retreat; this information will be updated annually as new images are obtained. In the Himalayan test case shown in Figure 3, commercial stereo imaging from SPOT was used to produce a DTM, and the results were then compared with a DTM generated with the USGS algorithm.

Different approaches can be used with a DTM to map the edges of glaciers in debris-covered areas. One approach developed at USGS involves automated drainage mapping. Another tool starts with a crude glacier centerline input by a person, and then the computer looks at a second derivative of the topography (slope changes) along transverse and longitudinal profiles starting from the centerline (Figure 3d). Satisfactory results were obtained in our test case with very little human involvement: 1) the command to run the program was typed; 2) a DTM image was selected; 3) a few points were clicked more or less down the apparent center of the glacier, and these were input to the algorithm, and then the outline appeared on the image a fraction of a second later.

This technique also worked for a glacier on Mount Rainier, a very different environment physiographically and topographically. The exact glacier outline is somewhat dependent on the spacing and placement of the clicked points, so the user repeated the process three times to get an edge that looked pretty good. This operator-dependency can be further reduced in the future. The glacier edge so obtained in the Nanga Parbat test case should coincide with the interior contact between the live ice and the prominent lateral moraine, which of course could be underlain by glacier ice.

The drainage mapping technique may produce output that more closely corresponds to the outer edge of the prominent lateral moraine. The ultimate test of any automated algorithm's output is comparison with field data. This has not yet been done, but it will be a key part of GLIMS.

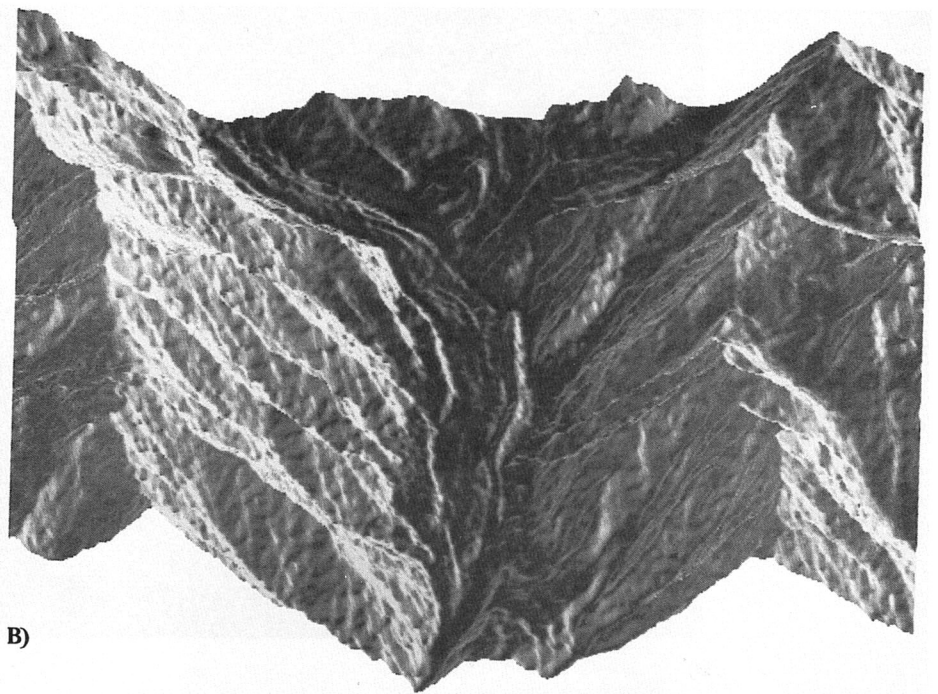
With combined use of multispectral and digital terrain model methods of mapping glaciers, it will be possible not only to map glaciers systematically worldwide, but to subdivide glaciers into major facies, such as debris-covered ice (for example, terminal ablation moraine and medial moraine), exposed glacier ice, and snow-covered ice. Although ASTER images generally will not be obtained exactly at the end of melt season, suitable semi-empirical statistical means may enable reasonably accurate extraction of the firm line elevation.

Automation, Model Validation

Increasing the automation of each image analysis step and providing for model validation and critical field verification of the glacier database are key requirements of GLIMS, to enable comprehensive and reliable analysis of the world's land ice using remote sensing.

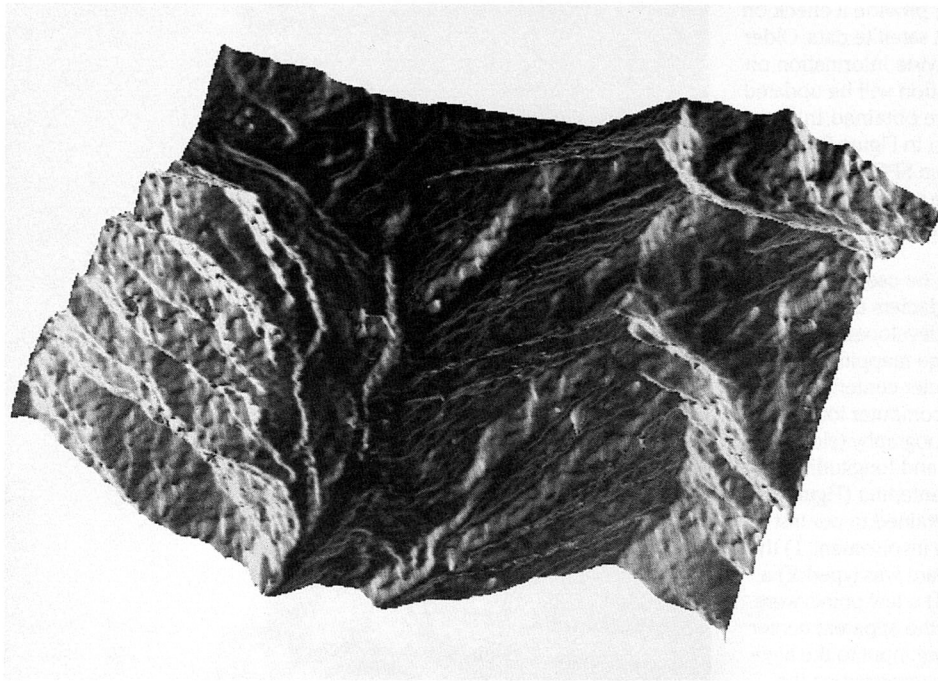


A)

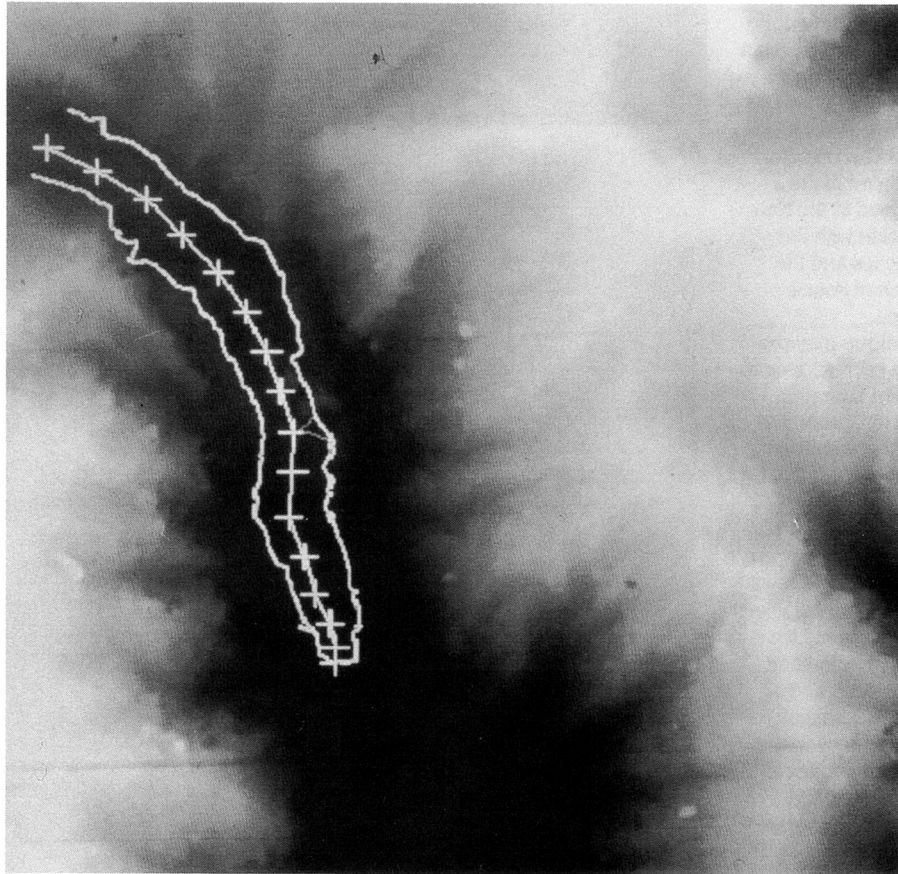


B)

Fig. 3. Glaciers in Nanga Parbat Range, Karakoram Himalaya. (a) Left SPOT image of a stereo pair of a 10 x 10 km portion of the Nanga Parbat Range, Himalaya. The long valley glacier is Raikot Glacier. (b) Shaded relief portrayal of a commercial DEM from the SPOT Corporation.



c)



D)

Fig. 3. (c) Shaded relief applied to a DTM produced at USGS, Flagstaff, using the SPOT image pair. Lateral moraines are very clearly visible. Total relief ranges from 2800 m elevation to 5600 m in this area. The moraines and shape of the glacial valley fill provide information on the glacier edge that, because of partial debris cover, probably would not be obtainable in this instance with multispectral classification. Multispectral data, however, can provide information about glacier facies (for example, snow, ice, and moraine). The perspective viewing geometry differs slightly between (b) and (c). For (c), no attempt has been made to filter or remove high-frequency artifacts that appear as small conical hills and subtle waves. The SPOT stereo pair and the commercial DEM were purchased from the SPOT Corporation and provided for this work courtesy of Jack Shroder and Michael Bishop of the Nanga Parbat project (NSF grant EAR 9418839) for purposes of GLIMS algorithm development and demonstrations. (d) Edge of Raikot Glacier determined by application of the slope-change algorithm (described in article) to the USGS DTM.

Algorithms are being written to minimize the need for user involvement.

GLIMS has been developed specifically with ASTER imaging in mind. However, the GLIMS consortium approach to satellite data analysis, the NSIDC glacier database, and modified GLIMS algorithms may be broadly usable with other satellite data. Use of Landsat-7 data is planned and will be needed especially in Antarctica. Use of radar satellite imaging (Radarsat, ERS-1, ERS-2, JERS, CRYOSAT, and Shuttle radar data) may provide the best means to measure the firm line. In interferometric mode, radar provides short-term change information, for example, glacier surface flow field.

CRYOSAT is a dedicated U.K.-led European Space Agency ice mission to be launched around 2003. Designed with two radar altimeters spaced 1 m apart, CRYOSAT will have tremendous capabilities to study glacier ice; it will provide more extensive absolute elevation and elevation change data complementing the data obtained by previous and concurrent multispectral and radar systems, including the data recently collected by the Shuttle Radar Topography Mission. CRYOSAT topography and synthetic aperture radar imaging will be especially vital where cloud cover limits ASTER, Landsat (expected in many notoriously cloudy glacierized mountain and Arctic island areas), and the Geoscience Laser Altimeter System (GLAS).

GLAS will be launched on the Ice, Cloud, and land Elevation Satellite (ICESat) around July 2001. Where clouds do not intervene, GLAS will provide absolute high-precision elevations and albedo data complementary to that obtained by radar and multispectral systems.

With planning well underway for a possible second and third Radarsat Antarctic Mapping Mission (AMM-2 and -3), tentatively scheduled for late 2000 and 2001 to acquire a highly accurate snapshot of ice motion for all Antarctic regions north of 80°S, a new NSIDC-led consortium is being launched, called VELMAP. This project will gather ice velocity data from all well-documented and accurate sources—current and historic, dating back to the early expeditions. NSIDC will build on its existing compilation of velocity data, viewable and downloadable at http://nsidc.colorado.edu/NSIDC/ANTARCT_VELOC/. The new data-

base will include latitude, longitude, speed, bearing, elevation, slope, and date range.

Two to Tango

Are glaciologists prepared to deal with the large volumes of data that have already begun to arrive on their desks? It has been common practice for a single pair of glacier scenes to consume an entire year's effort by a small team of glaciologists. The only way to take full advantage of new satellite capabilities and thousands of glacier images is to adopt a semi-automated approach to image analysis, such as that planned by GLIMS. But will measurements made by semiautomated image analysis be reliable? The only way to ensure that they are is to validate them with field glaciology data; this requires remote sensing and field glaciologists to interact heavily. It is just this sort of interaction and the systematic global approach to glacier measurements that the Terrestrial Observations Panel on Climate (TOPC) has in mind. TOPC, a U.N.-affiliated advisory panel, recently recommended that global glacier observation programs such as GLIMS and the Landsat-7 glacier program be adopted as demonstration projects for 21st-century global change monitoring.

It is not only technology and the scope of observations needed that challenge us. At a recent glaciology meeting, one glaciologist used a high-school dance analogy to describe the working relationship between field and remote sensing glaciologists. Boys and girls sit separately and gaze at each other with interest at the dance, but it takes some boldness and a new way of thinking to go beyond staring across the dance floor.

Data Inventory

For an inventory of existing glacier data, see <http://www.nsidc.colorado.edu/NSIDC/CATALOG/ENTRIES/GO1130.html>.

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